

The Effect of Submergence on the Distribution of Some Trace Elements within the Recent Sediments at Mosul Lake

Zeki. A. Aljubouri
Department of Geology
College of Science
University of Mosul

Hazim. A. Al-Kawaz
Research Center for Environment
and Water Resources
Mosul University

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ABSTRACT

Sixty five samples were collected from the eastern and western banks of Mosul lake at submergence levels of (305) and (327) (m.a.s.l) and analyzed for ten trace-elements (V, Co, Ni, Cu, Zn, Rb, Sr, Zr, Mo and Pb).

Four trends of variation in the concentrations of trace elements in the sediments of the lake and at both banks were observed. These are: (a) Increase at both (305) and (327)(m.a.s.l.) levels; the increase being more at the (305) level (V, Co, Zr, and Pb at the eastern bank). (b) Increase at both levels, the increase being less at (305) level (Zn). (c) Increase at both levels at equal or almost equal amounts, (Ni, Cu, Rb, Mo, Pb at the western bank) and (d) Decrease at both levels (Sr).

From the trends of variation, it may be envisaged that four factors were affecting the variation in concentrations of these elements within the sediments: dissolution of the relatively soluble mineral phases (carbonates and sulphates); mechanical removal of clay size particles (clay minerals); addition of organic matter and fourthly the slope of the banks.

تأثير الاغمار على توزيع العناصر الاثرية في الرواسب الحديثة في بحيرة الموصل

الملخص

تم جمع (65) نموذجاً من الرسوبيات الحديثة في الضفتين الشرقية والغربية لبحيرة الموصل ومن خطي اغمار (305) و (327) متراً فوق مستوى سطح البحر. وتم تحليل هذه النماذج لعشرة عناصر اثرية هي (V, Co, Ni, Cu, Zn, Rb, Sr, Zr, Mo and Pb).

وقد تم ملاحظة وجود اربعة انماط من التغيرات في تراكيز العناصر الاثرية في الرسوبيات الحديثة في البحيرة وهي: (أ) زيادة تراكيز العناصر في خطي اغمار (305) و (327) متراً فوق مستوى سطح البحر. والزيادة اكثر في خط (305) متر فوق مستوى سطح البحر. (العناصر Co, V, Zr, Pb في الضفة الشرقية فقط). (ب) زيادة في تراكيز العناصر في كلا خطي الاغمار. حيث تكون الزيادة اقل في خط الاغمار (305) متر مقارنة مع خط اغمار (327) (متر فوق مستوى سطح البحر). (ج)

زيادة في تراكيز العناصر بكميات متساوية او متساوية تقريباً (Pb , Mo , Rb , Cu , Ni) في الضفة الغربية) و (د) نقصان في كلا خطي الاغمار (Sr).
وقد تم تشخيص اربعة عوامل تؤثر على التغيرات في تراكيز هذه العناصر ضمن الرسوبيات الحديثة للبحيرة وهي : ذوبان الأطوار المعدنية ذات قابلية الذوبان العالية نسبياً (الكربونات والكبريتات) والازالة الميكانيكية للحبيبات الناعمة ذات الحجوم الطينية (المعادن الطينية) إضافة المادة العضوية إلى الرسوبيات و رابعاً درجة ميل الضفة والصخور المحيطة بها.

INTRODUCTION

Mosul lake is the largest reservoir lake on River Tigris. Its maximum storage capacity may exceed (13) billion cubic meters. The surrounding rocks belong to the Fat'ha (Middle Miocene) and Injana (Upper Miocene) Formations.

The Fat'ha Formation consists of repeated cycles of gypsum, marls and carbonates, whereas that of Injana Formation consists of sandstones, marls and conglomerates.

The rocks at the western bank of the lake have a very steep slope which often takes the form of a cliff, cutting through the rocks of Fat'ha Formation. On the other hand, the rocks and sediments of the eastern bank have a relatively gentle slope, not exceeding (30) degrees.

In April, 1987, the water of the lake reached a maximum level of (327) meters above sea-level (m.a.s.l) and stayed at this level for almost six months, then dropped to (315 m.a.s.l) for about (12) months. In November (1988), the level dropped to (303 m.a.s.l) and stayed at this level permanently.

The purpose of this work is to find the concentrations of ten trace elements (V, Co, Ni, Cu, Zn, Rb, Sr, Mo, Zr and Pb) within recent sediments of the previously submerged (305) and (327) levels at both eastern and western banks of the lake and to compare these concentrations with the unsubmerged part of the lake.

PREVIOUS STUDIES

Al-Naqib and Al-Taiee (1988) investigated the effects of transported sediments in the north-eastern part of the lake coming from Faida and Buqak valleys. They estimated that these transported sediments make up around (10%) of the total dead storage capacity of the lake.

Al-Rawi et al. (1990) investigated the pollution in Duhok valley and its effects on the quality of water in Mosul lake. Al-Kawaz (1991) studied the effects of erosion and the geochemistry and mineralogy of the recent sediments within (305) and (327) submergence levels of the lake. He concluded that the source of most mineral phases of the recent sediments of the lake is from the surrounding Fat'ha and Injana Formation.

METHODOLOGY

In November, 1988 and after the water level of the lake has dropped to (303 m.a.s.l) (96) samples representing the formerly submerged recent sediments at eastern and western banks of the lake from (305) and (327) levels were collected (Figure 1).

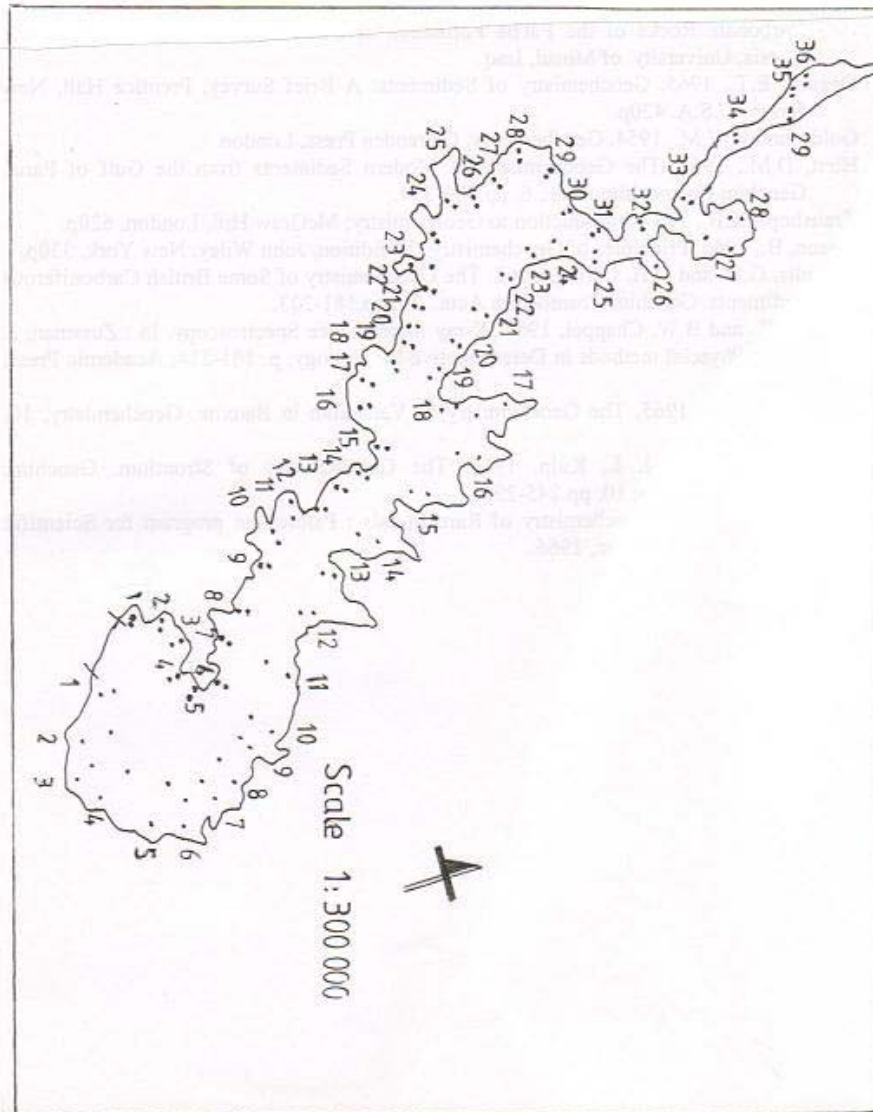


Fig. 1: Mosul lake, with sample location.

Of these, (65) samples were chosen for analyses, dried, and each sample was divided into three parts. One part was kept as a reference; the second and third were ground and used for mineralogical and geochemical analyses respectively. All analyses were carried out at the geochemical laboratories of the General Company for Geological Survey and Mining, Baghdad, Iraq.

Ten trace elements (V, Co, Ni, Cu, Zn, Rb, Sr, Mo, Zr and Pb) were analyzed using Phillips, XRF Spectrometer. The pellet method of Norrish and Chappel (1967) was applied. The results of analyses are shown in Tables (1,2,3 and 4). Table (5) shows the mean values for the concentrations of trace elements at each submergence level, as well as for the unsubmerged part of the sediments of the lake.

RESULTS and DISCUSSION

In order to understand the behavior of the studied trace elements during successive submergence of sediments; and in order to explain the variations in concentrations between different elements and variations of the same element at different submergence level, it is very important to know the geochemistry of these elements in sedimentary rocks. A summary of the main distribution of each element is given here after:

Vanadium (V) is mainly concentrated in the organic matter; in the residual fraction and in the clay minerals (Tenyakov, 1965).

Cobalt (Co), nickel (Ni) and copper (Cu) are mainly concentrated in the clay minerals, and in the heavy fraction including the sulphide phase. (Webber and Middleton 1961, Nicholls and Loring 1962, Degens 1965, Aljubouri et al. 1994); Zinc (Zn) is in dolomite; in clay minerals and in the sulphide phase (Mason, 1966; Aljubouri, 1972).

Rubidium (Rb) is concentrated in the clay minerals (illite); in K-feldspars and in the micas as well as the heavy fraction (Goldschmidt, 1954; Vlasov, 1967; Krauskopf, 1967). Strontium (Sr) is mainly found within the soluble phases (carbonates and sulphates), (Hirst, 1962; Turkian and Kulp, 1965; Aljubouri, 1972).

Zirconium (Zr) is found in the heavy fraction (Hirst, 1962; Al-Ubaidi, 1984). Molybdenum (Mo) and lead (Pb) are concentrated within the heavy fraction including the sulphide phase (Mason, 1966; Krauskopf, 1956; Aljubouri and Sulayman, 1996).

Al-Kawaz (1991) has found that the mineralogy of the submerged and unsubmerged recent sediments of Mosul lake consists mainly of carbonates (dolomite and calcite), quartz and free iron oxides (mainly hematite, Fe_2O_3) together with the clay minerals (illite, chlorite, kaolinite and motmorillonite).

Because of similarities in mineral phases between recent sediments and the surrounding rocks of the lake. Al-Kawaz (1991) concluded that the source rocks are mainly from the surrounding Fat'ha Formation (Middle Miocene) and to a lesser extent from Injana Formation (Upper Miocene).

It may be envisaged here that the submerged sediments of the lake are subjected to four factors ; these are: dissolution of the soluble mineral phases (calcite and gypsum, then dolomite to a lesser extent); mechanical removal of clay-size particles from sediments and hence accumulation of heavy fraction minerals. Accumulation of organic matter, due to the decay of algae , leading to the formation of organo-metallic complexes, and fourthly the slope of the surrounding rocks and sediments . The lower the slope of the sediments, the more area is exposed to submergence.

Table 1: Trace Element Analyses (ppm) at Eastern bank, 305 Level (305EA), (n=15).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
EA1	142	11	195	28	53	37	155	166	6	7
EA2	163	8	50	29	26	24	193	78	5	15
EA3	180	7	66	25	10	20	231	51	5	6
EA4	145	8	89	23	22	35	221	83	5	7
EA5	165	9	208	32	36	36	319	100	5	11
EA6	220	27	290	29	64	50	210	136	7	7
EA7	180	21	320	27	72	61	214	168	14	40
EA8	200	26	230	32	64	48	231	136	7	17
EA9	210	21	280	36	74	53	214	143	14	49
EA10	350	8	340	29	74	41	295	92	7	5
EA11	124	5	110	28	23	64	167	191	6	5
EA12	220	7	160	33	34	41	284	134	5	32
EA13	270	7	280	31	43	27	1013	98	5	68
EA14	220	5	200	41	30	29	193	30	5	23
EA15	200	12	250	47	75	51	140	75	5	56

Table 2: Trace Element Analyses (ppm) at Eastern Bank, 327 Level, (327EB) of the Lake (n=14).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
EB16	230	7	205	31	67	52	171	135	5	14
EB 17	210	7	152	19	13	10	271	47	5	5
EB 18	210	9	83	25	18	10	229	46	5	15
EB 19	119	7	205	27	34	33	240	87	5	5
EB 20	104	21	290	31	104	44	198	137	6	19
EB 21	192	9	160	28	40	27	190	94	5	12
EB 22	270	17	55	37	40	55	193	118	7	15
EB 23	186	8	260	29	72	54	217	154	5	11
EB 24	240	8	430	37	92	66	519	105	5	21
EB 25	240	11	180	34	44	38	878	110	6	40
EB 26	140	5	32	19	15	19	221	72	5	6
EB 27	145	7	230	24	60	35	176	131	6	5
EB 28	170	7	310	36	43	27	521	94	5	41
EB 29	120	5	250	28	75	108	220	75	5	5

Table 3: Trace Element Analyses (ppm) at Western Bank, 305 Level (305 WA) of the Lake (n = 22).

S.N.	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
WA1	5	15	140	15	11	10	170	11	5	5
WA2	240	18	330	37	56	49	170	56	7	9
WA3	250	15	350	29	48	30	220	57	5	40
WA4	220	18	320	32	52	44	90	106	7	25
WA5	145	15	250	25	51	43	220	120	7	6
WA6	41	29	145	23	26	21	740	76	7	8
WA7	58	15	154	29	16	10	480	68	5	5
WA8	58	25	144	22	26	18	740	68	5	5
WA9	78	15	180	22	18	17	170	75	7	5
WA10	380	16	250	19	53	34	170	118	6	21
WA11	80	15	162	28	23	25	130	94	6	5
WA12	48	15	160	21	19	18	100	81	7	5
WA13	58	25	100	26	20	18	170	54	5	7
WA14	7	29	64	26	13	10	480	53	7	5
WA15	34	20	36	21	16	11	130	80	5	5
WA16	11	21	200	20	22	24	220	80	7	5
WA17	19	15	40	32	20	27	100	60	7	9
WA18	100	16	110	21	36	29	260	113	5	5
WA19	110	17	110	19	32	35	864	195	5	6
WA20	350	30	350	70	106	64	130	137	7	25
WA21	350	48	360	62	115	73	260	164	7	31
WA22	340	32	340	62	116	76	191	138	8	20

Table 4: Trace Element Analyses (ppm) at Western Bank, 327 Level (327 WB) of the Lake (n = 14).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
WB23	5	5	5	16	10	10	130	10	5	5
WB24	390	11	360	41	93	51	170	93	6	8
WB25	165	8	192	25	47	43	570	105	5	13
WB26	142	12	183	18	20	17	170	62	6	5
WB27	40	17	41	29	12	17	190	58	5	5
WB28	36	16	98	25	18	10	200	67	5	5
WB29	140	22	130	29	36	25	220	72	5	16
WB30	14	16	80	28	20	10	420	57	5	5
WB31	7	15	25	24	12	22	188	55	5	5
WB32	11	18	44	29	21	30	153	112	5	5
WB33	22	19	260	20	70	59	140	118	5	5
WB34	220	16	90	26	42	32	1300	115	5	24
WB35	280	18	305	50	106	72	170	140	6	31
WB36	300	20	345	34	120	83	189	141	5	36

Examination of Table (5) reveals that vanadium (V) content at the eastern bank of the lake is higher at (305) level than that at (327) level, this is because the (305) level has remained submerged for longer period (18) months than the (327) level (6

months). This means that more organic matter has accumulated at the (305) level and more removal of clay-size particles and the concentration of residual fraction. These processes have lead to more concentration of vanadium at the (305) level, compared with (327) level.

Table 5: Mean Values for the Analyses of Trace Elements (ppm) Within (305) and (327) (m.a.s.l.) Submergence Levels of the Recent Sediments of the Lake Compared with Unsubmerged Parts.

Sediment	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
EA	199	12	205	31	47	41	272	112	6	23
EB	184	9	203	30	51	41	289	100	5	15
*EC	140	7	152	20	29	23	303	80	N.D.	6
WA	136	21	195	30	41	31	282	91	6	12
WB	127	15	154	28	45	34	301	86	5	12
*WC	100	9	150	24	35	24	345	70	N.D.	7

EA: 305 level, Eastern Bank

EB: 327 level, Eastern Bank

EC: Unsubmerged Part, Eastern Bank

WA: 305 level, Western Bank

WB: 327 level, Western Bank; WC: Unsubmerged Part, Western Bank

N.D.: Not Detected

*: From Al-Kawaz, 1991

Similar situation exist at the western bank at both (305) and (327) levels, but to a lesser extent (Tables 5 and 6) due to the steeper slope at this bank, which means lesser area is exposed to submergence. The steep slope may cause the sliding of large masses of rocks into the water, leading to a continual exposure of fresh surfaces under the effects of the aforementioned factors, and hence a relative decreases in their effects.

Because of the rather low concentrations of (Co), (Mo) and (Pb) at eastern bank), the discussion of their behavior under submergence is not very reliable.

The trace elements (Co) and (Ni) at the eastern bank; (Cu), (Rb) and (Mo) at both banks and (Pb) at the western bank, show an increase in their concentrations due to submergence relative to the unsubmerged part (Table 5). However, the increase at (305) level either equal to (Rb and Pb) or slightly more than at (327) level.

Table 6: Percentage Increase or Decrease in Trace Elements at the Four Submerged Levels Relative to the Unsubmerged Part (in Table 5).

Element	305 E (A)	327 E (B)	305 W(A)	327 W(B)
V	42.1%	31.4%	36.0%	27.0%
Co	71.4	28.6	133.3	66.7
Ni	34.9	33.6	30.0	2.7
Cu	10.7	7.1	25.0	16.7
Zn	62.1	75.9	17.0	28.6
Rb	78.3	78.3	29.2	41.7
Sr	-10.2	-4.6	-18.3	-12.8
Zr	40.4	25.0	30	23
Mo	X6(600%)	X5	X6	X5
Pb	283	150	71.4	71.4

It is possible that each of these elements is present in at least two mineral phases (clay minerals and heavy fraction), in almost equal amounts. The decrease in their concentrations due to the removal of clay size particles is counter balanced by the increase in concentration of the heavy fraction, including the sulphide phase.

At the western bank and at (305) level, (Ni) shows a marked increase (195ppm) among these elements compared with the (327) level (154 ppm) and with unsubmerged part (150 ppm). It is possible that (Ni) is concentrated within the sulphide phase more than within the clay minerals which are continually removed due to successive and longer submergence of (305) level relative to (327) level.

Zinc (Zn) unlike the rest of elements shows a decrease in its concentration at (305) levels of both banks. This is possibly due to its presence in the carbonate phase (dolomite), which has been subjected to dissolution at (305) level more than at (327) level due to longer submergence of the (305) level.

Strontium (Sr) is the only element that shows a marked decrease in its concentration at both levels and at both banks and the decrease is more at (305) level (Tables 5 and 6). It is clear that the decrease in (Sr) is due to the dissolution of the relatively soluble phases of carbonates (calcite and dolomite) and sulphates (gypsum), where (Sr) substitutes for calcium (Ca^{2+}) in these phases. The longer submergence period of the (305) level causes more (Sr) to be dissolved at this level.

Zirconium (Zr) exhibits a marked increase at both (305) and (327) levels and at both eastern and western banks. The increase is more at (305) level (Table 5). It is well known that zirconium in sedimentary rocks is concentrated in the heavy mineral zircon ($ZrSiO_4$). The longer the submergence period, the more fine grained fraction is removed, which means the more is the concentration of the heavy fraction

The percentage increase of (Zr) at (305) level at both banks (east and west (40% east), and (30% west) is more than that at (327) level (25% east) and (23% west). This is due to longer submergence period at the (305) level than the (327) level.

At the eastern bank, the element lead (Pb) shows an increase at both (305) and (327) levels being more at the (305) level (Table 5). This means that (Pb) is mainly concentrated within the heavy fraction, the sulphide phase at this bank.

From the foregoing discussion, it seems that there are four patterns of variation in the concentrations of the studied trace elements in the recent sediments of Mosul lake; these are: (a) Increase in their contents at both (305) and (327) levels; the increase being more at the (305) level (V, Co and Pb at the eastern bank only) (b) Increase at both (305) and (327) levels; the increase being less at (305) level (Zn) (c) Increase at both levels at equal or almost equal amounts (Ni, Cu, Rb, Mo and Pb at the western bank) and (d) decrease at both levels (Sr).

Inter-element correlation matrices for (305) and (327) levels at both eastern and western banks (Tables 7, 8, 9 and 10) confirms the pattern of distributions of the ten trace elements discussed above.

The elements that show an increase in their concentrations at the (305) or (327) level exhibit positive correlations between them. For example, at the eastern bank and at (305) level (Table 7), vanadium (V) positively correlates with (Co, Ni, Cu, Zn, Rb, Zr, Mo, and Pb). All these elements exhibit an increase in their concentrations at (305) level, compared with the unsubmerged part of the sediment (Table 5).

On the other hand, these elements show negative correlation with (Sr) which shows decrease in its concentration at the (305) level (and at 327 level). Similar relationship hold for most of these trace elements in the other three tables (Tables 8, 9 and 10).

Table 7: Correlation Coefficient Matrix for the Trace Elements, 305 level, Eastern Bank (n = 15).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
V	1.00									
Co	0.43	1.00								
Ni	0.82	0.48	1.00							
Cu	0.73	0.03	0.66	1.00						
Zn	0.90	0.60	0.87	0.74	1.00					
Rb	0.72	-0.2	0.51	0.01	0.54	1.00				
Sr	-0.20	-0.35	-0.25	-0.50	-0.26	-0.58	1.00			
Zr	0.49	0.46	0.24	0.01	0.34	0.84	-0.11	1.00		
Mo	0.47	0.54	0.44	0.61	0.45	0.61	-0.13	0.46	1.00	
Pb	0.49	0.03	0.72	0.01	0.54	0.01	0.65	0.01	-0.02	1.00

Table 8: Correlation Coefficient Matrix for the Trace Elements, 327 level, Eastern Bank (n = 14).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
V	1.00									
Co	0.05	1.00								
Ni	0.70	0.11	1.00							
Cu	0.56	0.09	0.75	1.00						
Zn	0.47	0.36	0.79	0.70	1.00					
Rb	0.39	0.0	0.51	0.73	0.69	1.00				
Sr	-0.52	-0.31	-0.54	-0.29	-0.26	-0.23	1.00			
Zr	0.64	-0.02	0.26	0.07	0.01	-0.02	0.13	1.00		
Mo	0.13	0.73	-0.06	-0.10	0.13	0.10	0.09	0.07	1.00	
Pb	0.46	0.31	0.67	0.83	0.40	0.49	-0.18	0.22	0.22	1.00

Table 9: Correlation Coefficient Matrix for the Trace Elements, 305 Level, Western Bank (n = 22).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
V	1.00									
Co	0.33	1.00								
Ni	0.92	0.36	1.00							
Cu	0.72	0.63	0.78	1.00						
Zn	0.89	0.59	0.86	0.89	1.00					
Rb	0.30	-0.12	0.44	0.17	0.16	1.00				
Sr	-0.26	-0.45	-0.33	-0.28	-0.23	-0.10	1.00			
Zr	0.63	0.44	0.50	0.52	0.70	-0.11	-0.02	1.00		
Mo	0.51	0.52	0.46	0.67	0.71	-0.01	0.07	0.45	1.00	
Pb	0.80	0.30	0.85	0.06	0.70	0.68	-0.26	0.35	0.33	1.00

Table 10: Correlation Coefficient Matrix for the Trace Elements, 327 Level, Western Bank (n = 14).

	V	Co	Ni	Cu	Zn	Rb	Sr	Zr	Mo	Pb
V	1.00									
Co	0.92	1.00								
Ni	0.71	0.53	1.00							
Cu	0.92	0.49	0.68	1.00						
Zn	0.81	0.65	0.89	0.82	1.00					
Rb	0.68	0.67	0.83	0.72	0.95	1.00				
Sr	-0.23	-0.19	0.00	0.13	-0.01	-0.02	1.00			
Zr	0.58	0.72	0.69	0.65	0.76	0.86	0.23	1.00		
Mo	0.62	0.39	0.45	0.52	0.62	0.57	-0.18	0.36	1.00	
Pb	0.68	0.83	0.49	0.72	0.75	0.75	0.34	0.70	0.53	1.00

CONCLUSIONS

1. There are four trends of variation in the concentrations of the studied ten trace elements, in the submerged sediments relative to the unsubmerged; these are: (a) Increase at both (305) and (327) levels; the increase being more at the (305) level (V, Co, Zr and Pb); (b) Increase at both (305) and (327) levels. The increase being less at (305) level (Zn). (c) Increase at both (305) and (327) levels at equal or almost equal amounts (Ni, Cu, Rb, Mo and Pb at the western bank) and (d) decrease at both levels (Sr).
2. There are four factors affecting the distribution of trace elements within the submerged sediments of Mosul lake; these are: dissolution of relatively soluble mineral phases (carbonates and sulphates); mechanical removal of clay-size particles; addition of organic matter; and the slope of the submerged sediments and surrounding rocks.

REFERENCES

- Aljubouri, Z.A.J., 1972. Geochemistry, Origin and Diagenesis of Some Triassic Gypsum Deposits and Associated Sediments in the East Midlands, England. Unpublished Ph.D. Thesis, University of Nottingham, England, U.K.
- Aljubouri, Z.A.J., Al-Kattan, M. and Al-Sayegh, A.Y. 1994. Geochemistry, Mineralogy and Colour of Gercus Formation at Zawita, Iraq Africa Geoscience Review, 1(4), pp.547-556.
- Aljubouri, Z.A.J. and M.D. Sulayman, 1996. Mineralogy and Geochemistry of Gypsum Rocks of Al-Fath'a Formation at West Butma Area, Northern Iraq. Raf. Jour. Sci., 7(1), pp.114-128.
- Al-Kawaz, H.A.M., 1991. The effects of Erosion on the Geochemical and Mineralogical Variations of the Recent Sediments, within Submerged Part of Mosul Lake: Confidential. Mosul Research Centre for Dams and Water Resources. University of Mosul, Iraq. (In Arabic).
- Al-Naqib, S. Q. and Th. M. Al-Taiee, 1988. The Effect of Suspended Sediment load, Transported from Fayda and Buqak Wadies, on Mosul Lake as related to Watershed Characteristics. Confidential Mosul Dam Research Centre, University of Mosul, Iraq.

- Al-Rawi, S.M., Mustafa, M.M. and Al-Kawaz, H.A.M., 1990. A Study of the Pollution in Duhok Valley and its Impact Upon Mosul Lake Water Quality. Mosul Dam Research Centre. Mosul University, Mosul, Iraq. Confidential.
- Al-Ubaidi, S.A., 1984. Geochemistry, Mineralogy and Petrography of the Argillaceous Carbonate Rocks of the Fat'ha Formation at Butma West Area. Unpubl. M.Sc. Thesis, University of Mosul, Iraq.
- Degens, E.T., 1965. *Geochemistry of Sediments: A Brief Survey*; Prentice Hall, New Jersey, U.S.A. 420p.
- Goldschmidt, V.M., 1954. *Geochemistry*, Clarendon Press, London.
- Hirst, D.M., 1962. The Geochemistry of Modern Sediments from the Gulf of Paria. *Geochim Cosmochim Acta.*, 6, pp.309-334.
- Kraushopf, K.B., 1967. *Introduction to Geochemistry*; McGraw Hill, London, 620p.
- Mason, B., 1966: *Principles of Geochemistry*; 3rd Edition, John Wiley, New York, 330p.
- Nicholls, G.D. and D.H. Loring, 1962. The Geochemistry of Some British Carboniferous Sediments. *Geochim Cosmochim Acta.*, 26, pp.181-223.
- Norrish, K. and B.W. Chappel, 1967. X-ray fluorescence Spectroscopy. In : Zussman, J. (Ed) : *Physcial methods in Determinative Mineralogy*, p. 161-214; Academic Press, London.
- Tenyakov, V.A., 1965. The Geochemistry of Vanadium in Bauxite. *Geochemistry*, 10, 553.
- Turkian, K.K. and J. L. Kulp, 1956. The Geochemistry of Strontium. *Geochim. Cosmochim. Acta*; 10, pp.245-296.
- Vlasov, K.A., 1964. *Geochemistry of Rare Metals : Palestinian program for Scientific Translation*; Jerusalem, 1966.